

Thelma began as a disturbance in the monsoon trough approximately 100 nm (185 km) north-northeast of Saipan. Although Thelma's forecast track posed few problems, analysis of the cyclone was not straight-forward due to Thelma's very abnormal structure.

Satellite imagery began showing intensification of convective activity in the maximum cloud zone within the latitude belt 10-20N from southeast Asia to the Marshall Islands on 23 September. A significant area of convection developed in the monsoon trough near Saipan early on 23 September. This convection gradually increased and began suggesting the presence of a surface circulation later on the same day. Data at 500 mb, however, indicated that the curvature in the cloud signature was associated with a mid-tropospheric circulation as no circulation was apparent from surface/gradient wind data. circulation at the surface/gradient level was finally analyzed just south of Guamat 2412002. By 2500002, the low-level circulation that produced Thelma was analyzed over the Northern Mariana Islands (Fig. 3-22-1). This circulation continued to develop and drift westward while becoming the dominant circulation in that portion of the monsoon trough.

From her onset, Thelma did not display classical tropical storm characteristics. Height gradients observed by reconnaissance aircraft at 700 mb were very flat; thus maximum winds near the center were significantly lower than suggested by the central sea-level

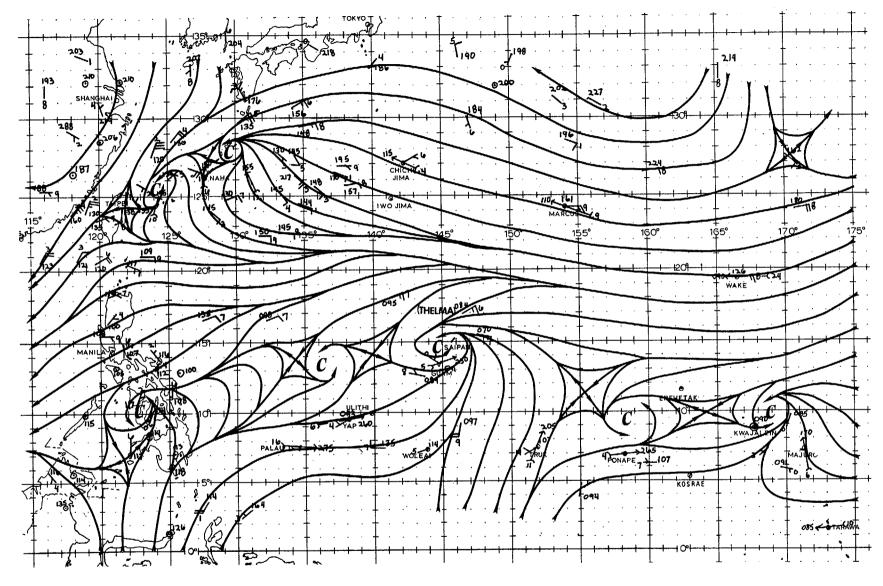
pressure (Atkinson and Holliday, 1977). Also, the maximum wind band was some distance away from the center, and the 700 mb temperature field showed higher temperatures outside the cyclone center for most of the early aircraft penetrations. Table 3-3 presents a summation of aircraft data and highlights the points presented above. Figure 3-22-2 shows the reconnaissance data plot for the last daylight penetration of Thelma. The wind field and other data presented on the plot are fairly representative of Thelma's entire life.

As previously stated, Thelma's track presented no real problems. Streamline analyses at 500 mb showed that Thelma developed just south of a break in the subtropical ridge. After following a northwesterly course, Thelma first turned northward and almost immediately thereafter began to track northeastward. Cyclogenesis/frontogenesis occurred simultaneously in a baroclinic zone that persisted throughout Thelma's life in the area from Okinawa northeastward to a point off the coast of Japan. The continual presence of this surface trough appears to be one of the factors that directed Thelma's northeastward movement (Fig. 3-22-3). Upper-level steering was provided by relatively strong westerlies which reached south to the Bonin and Volcano Islands. During her northeast trek, Thelma reached maximum intensity of 55 kt (28 m/sec). Further intensification was probably sup-pressed by restrictions on her upper-level outflow. Thelma continued to accelerate toward the northeast and transitioned into an extratropical low pressure system by 04002 on 30 September.

TABLE 3-3

Date/Time	Maximum Temperature		Central Pressure	Intensity (KT)	Bearing/Range
	Inside Center <sup>O</sup> C	Outside Center OC	MB	Observed Atkinson/ Surface Holliday	of MAX FLT LVL Wind (DEG/NM)
27/1530Z	12	13	993	N/O 42	100/136
28/0308Z	14	15	989*	30 48	320/110
28/14182	12	15	987	N/O 50	240/40
29/025 <b>8</b> Z	14	10	982	55 57	350/110
29/1500Z	13	13	981*	N/O 59	120/150

Aircraft data extracted from detailed vortex messages and peripheral data observations. (Asterisks indicate central pressure extrapolated from 700 mb data.)



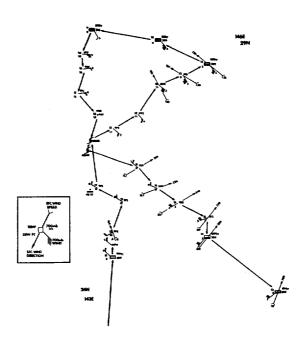


FIGURE 3-22-2. Plot of aircraft reconnaissance data for the 2903002 September 1980 fix of Tropical Storm Thelma.

